

Performance of the site-adapted CAMS database and locally adjusted cloud index models for estimating global solar horizontal irradiation over the Pampa Húmeda

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Abstract

CAMS provides global solar radiation estimates for clear-sky (McClear model) and all-sky (Heliosat-4 method) conditions, the latter based on MSG satellite information. A performance assessment of these estimates (with site-adaptation and spatial smoothing) is done, using hourly data from 10 sites in the Pampa Húmeda region of South America. Two locally adjusted Cloud Index Models (CIM) using GOES-East satellite information are also evaluated. One of them (CIM-ESRA) is based on the ESRA clear-sky model and the other (CIM-McClear) on the McClear clear-sky model. Under clear-sky conditions, the site-adapted McClear is found to perform best with a relative root mean square deviation (rRMSD) of 2.8%. However, in the presence of clouds in the real atmosphere, the model tends to provide lower clear-sky estimates than the ESRA model which, in our implementation, is only sensitive to average atmospheric trends. Under all-sky conditions, both CIMs show a small but consistent underestimation of -1.1% in the region and perform significantly better than the site-adapted Heliosat-4, with rRMSDs of 12.1% (CIM-McClear), 12.5% (CIM-ESRA) and 16.8% (site-adapted Heliosat-4). This performance difference is not a statement about the relative quality of the models, since it can be explained by the difference in satellite view angle (significantly higher for the MSG satellite than for the GOES-East satellite). The performance downgrade due to using MSG satellite images out of their recommended area is quantified. Both CIMs, based on using GOES-East imagery, provide accurate solar irradiation estimates over this region and can be extended to other areas of Latin America.

Keywords: Solar resource assessment, GHI, CAMS, satellite view angle, GOES satellite, hybrid models.

1. Introduction

The uncertainty of solar resource assessment is one of the main factors affecting the financial risk evaluation of large scale solar energy projects. This assessment ideally requires long-term, controlled quality, solar irradiation ground data for the project's site. Since this information is not usually available for a given project location, irradiation estimates based on geostationary satellite images are frequently used. These images provide the temporal and spatial resolution required for modeling a highly variable phenomena like ground level solar irradiation. The general idea is to quantify cloudiness using satellite information and use it to attenuate the clear-sky irradiation. Different models exist for this purpose (Perez et al., 2002; Ceballos et al., 2004; Rigollier et al., 2004; Cebecauer et al., 2010; Alonso-Suárez et al., 2012; Qu et al., 2017).

This work focuses on models for estimating ground-level solar global horizontal irradiation (GHI) from satellite information, working at the hourly time scale. Typical biases for hourly GHI satellite-derived estimates are within $\pm 3.5\%$ of the ground measurement's average (Perez et al.,

2013), excluding special cases such as tropical regions, polluted areas, high latitude areas with snow, mountains or complex island sites, where higher biases can occur. Typical dispersion for hourly estimates (as quantified by the relative root mean square deviation or RMSD) for arid and semi-arid climates is in the range 7-20% and, for areas with more complex cloud dynamics, between 15-30% (Perez et al., 2013). Uncertainty can be reduced by spatial smoothing or site-adaptation techniques, the latter by post-processing the estimates using good-quality ground measurements (Polo et al., 2016).

Models for solar satellite-based estimation can be classified as empirical (Tarpley, 1979; Justus et al., 1986; Cano et al., 1986; Alonso-Suárez et al., 2012), physical (Ceballos et al., 2004; Qu et al., 2017) and hybrid (or semi-empirical) models (Perez et al., 2002; Rigollier et al., 2004; Cebecauer et al., 2010). Empirical models rely on parametrizations between solar irradiation and other variables (i.e. satellite-derived cloudiness, solar zenith angle) with a set of parameters that are adjusted from ground measurements. Physical models attempt to model in detail the radiative transfer of solar irradiance through the atmosphere. The Heliosat-4 method (Qu et al., 2017) is a recent example of a successful physical model based on Meteosat Second Generation (MSG) satellite images and radiative transfer

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47 calculations. Hybrid models have an underlying physical
 48 structure with a few adjustable parameters. Both phys-
 49 ical and hybrid models are potentially accurate provided
 50 the required information is available with sufficient qual-
 51 ity. However, this information (i.e. aerosol optical depth,
 52 water vapor content, cloud type and phase, among oth-
 53 ers) is not always available with sufficient accuracy and
 54 spatial/temporal resolution. On the other hand, empiri-
 55 cal models require high quality ground measurements of
 56 adequate length to adjust their parameters and their es-
 57 timates cannot be extrapolated to other regions. Hybrid
 58 models provide a trade-off between empirical and phys-
 59 ical models. A common hybrid model approach is to use a
 60 physical clear-sky model modulated by a satellite-derived
 61 cloud index to generate solar irradiation estimates under
 62 all-sky conditions. These models are collectively known as
 63 CIM (Cloud Index Methods). The SUNY model (Perez
 64 et al., 2002) and the early Heliosat models, Heliosat-1
 65 (Beyer et al., 1996) and Heliosat-2 (Rigollier et al., 2004)
 66 are well known examples of this kind.

67 Another satellite-based model (of the empirical type)
 68 named BD-JPT, as it evolved from an original formulation
 69 by Justus, Paris and Tarpley (Justus et al., 1986), has been
 70 recently evaluated for the same region considered in this
 71 work (Alonso-Suárez et al., 2012). This model has been
 72 locally adjusted to ground data and used as a basis for the
 73 solar resource distribution map in Uruguay (Alonso-Suárez
 74 et al., 2014). In Alonso-Suárez et al. (2012), both the origi-
 75 nal JPT model and the improved brightness-dependent
 76 version (BD-JPT) have been evaluated for this region and
 77 showed interesting results: a relative RMSD of 13% at the
 78 hourly level with negligible bias was found. These mod-
 79 els were implemented with the same GOES-East satellite
 80 information used in this work for locally adjusted CIMs,
 81 in particular, using the same spatial averaging procedure
 82 described in Subsection 3.2.4.

83 This work provides, among other contributions sum-
 84 marized at the end of this Section, a first representative
 85 assessment for the Pampa Húmeda region (South East part
 86 of South America) of the Heliosat-4 method (Qu et al.,
 87 2017). The clear-sky part of this method, known as the
 88 McClear model (Lefèvre et al., 2013), is based on a param-
 89 etrization of the libRadtran libraries output (Mayer & Kylling,
 90 2005). McClear uses atmospheric information from the
 91 CAMS (Copernicus Atmosphere Monitoring Service) and
 92 ground albedo from the sun-synchronous orbiting MODIS
 93 satellite (Moderate Resolution Imaging Spectroradiome-
 94 ter) to estimate clear-sky irradiation. These clear-sky es-
 95 timates are combined with the McCloud model to produce
 96 the Heliosat-4 all-sky irradiation estimates. Cloud infor-
 97 mation and properties are derived from multiple spectral
 98 channels of the MSG satellite using the APOLLO/SEV
 99 algorithm (WDC, 2015). Here, the performance of this
 100 method based on the MSG satellite is compared against
 101 two locally adjusted hybrid CIMs which use cloud infor-
 102 mation derived from GOES-East satellite images. One
 103 of these CIMs is based on the ESRA (European Solar

Radiation Atlas) clear-sky model (Rigollier et al., 2004)
 and the other is based on the McClear model. The per-
 formance comparison between a spatially smoothed site-
 adapted model based on MSG images and two locally-
 adjusted models based on GOES-East images helps to
 quantify the impact of the different viewing angles with
 which both satellites see the area of interest and empha-
 sizes the importance of selecting the most adequate satel-
 lite information for each region.

Clear-sky models are important as a basis for CIMs
 and can be used to provide reliable upper bounds for au-
 tomated quality assessment of ground data or to com-
 pute the clear-sky index required by several applications,
 such as variability assessment or solar resource forecast-
 ing. Therefore, the performance of the two clear-sky mod-
 els used in this work (ESRA and McClear) is also eval-
 uated. These models differ markedly in their description
 of the atmosphere. While McClear captures the daily and
 intra-day atmospheric variability, ESRA has a single pa-
 rameter which, in our implementation, describes average
 atmospheric information, as explained in Subsection 3.3.
 This allow us to show a significant difference in the be-
 havior of the clear-sky estimates whether the actual real
 atmosphere is clear-sky or cloudy.

The main contributions of this work can be summa-
 rized as follows:

- Compares the performance of satellite-based models
 for all-sky hourly irradiation estimate based on dif-
 ferent geostationary satellite information and quan-
 tifies the impact of using satellite estimates out of
 their recommended area (i.e. satellite zenith angle
 larger than 60°).
- Compares two clear-sky models that differ in their
 capability for modelling the short-term atmospheric
 variability, in particular, by using water vapour as an
 input. For instance, it is found that when clouds are
 present in the atmosphere, modelling the short-term
 variability provides lower clear-sky estimates than
 using average atmospheric information. The ratio-
 nale is that the presence of clouds correlates with
 higher water vapour contents in the atmosphere and
 this results in lower clear-sky estimates.
- Provides a first representative performance assess-
 ment of the Heliosat-4 method and locally adjusted
 CIMs for the Pampa Húmeda area, including the
 gain quantification of a simple site-adaptation pro-
 cedure applied to the Heliosat-4 estimates.
- Quantifies the effect of the satellite information spa-
 tial smoothing in the region to reduce the uncer-
 tainty of hourly GHI estimates.

The article is organized as follows: Section 2 describes
 the satellite images, the ground data and the CAMS prod-
 ucts used in this work, including a short discussion on the

157 typical view angles from each satellite. In Section 3 the locally implemented models and their local-adaptations are
 158 discussed. This Section also describes the spatial smoothing procedure (applied to ensure that both satellite data
 159 sets have the same spatial averaging required for a fair comparison). In Section 4 the performance assessment of
 160 these models is done and discussed. Finally, our conclusions are summarized in Section 5.
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165 2. Data

166 The area of interest in this work is the part of south-eastern South America known as Pampa Húmeda, within
 167 latitudes 28°S and 36°S. As shown in Figure 1, it includes all the territory of Uruguay and parts of Argentina and
 168 southern Brazil. It is geographically homogeneous (mostly plain grasslands) with temperate climate and no important
 169 elevations. Although temperatures in winter can drop a few degrees below freezing point, snow episodes are rare.
 170 It is classified in the updated Köppen-Geiger climate classification (Peel et al., 2007) mostly as Cfa (temperate, with-
 171 out dry season, hot summers) with the exception of two small coastal regions dominated by the influence of the
 172 Atlantic Ocean and classified as Cfb (temperate, without dry season, warm summers).
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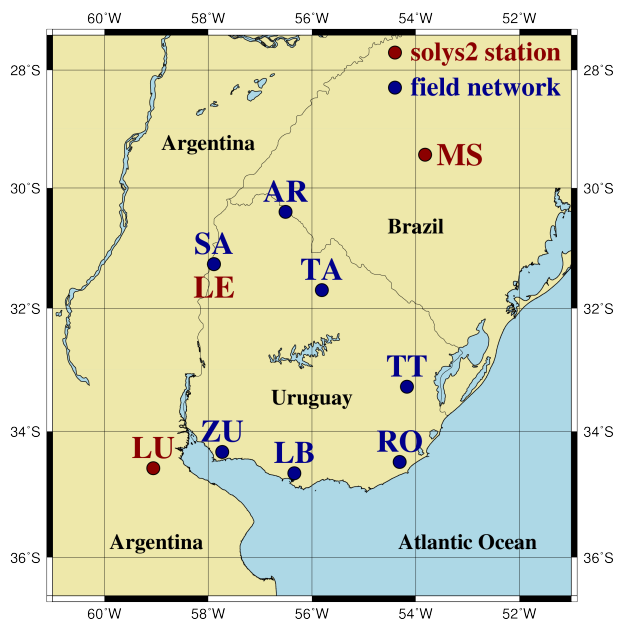


Figure 1: Location of the ground measurements stations.

180 2.1. Ground measurements

181 Ten series of GHI ground measurements are considered in this work. They belong to two groups, based on
 182 the quality of the instruments and the declared maintenance schedule at each site. The first group is composed
 183 by three ground stations located in Uruguay, Argentina and Brazil, whose equipment and procedures comply with
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BSRN requirements (McArthur, 2005): (i) the Solar Energy Laboratory experimental research facility (LE) in the
 north-western part of Uruguay (LES, <http://les.edu.uy>), (ii) the São Martinho da Serra station (MS), formally a
 BSRN site, and (iii) the Luján station (LU) located 50 km from Buenos Aires (Argentina) at a specialized research
 laboratory of the Luján National University (GERSolar, <http://www.gersol.unlu.edu.ar/>). At these sites, GHI is
 measured with ventilated secondary standard pyranometers and direct and diffuse irradiance are measured using
 precision solar trackers. Data are recorded as 1-min averages of several measurements. The LE instruments are
 calibrated every two years against a secondary standard (Kipp & Zonen CMP22) kept in storage and with traceability
 to the World Radiometric Reference (WRR). At the LU site, instruments are compared periodically against a Kendall
 absolute cavity radiometer, calibrated in 2018 with traceability to the WRR, which is stored and used sporadically
 as a reference. The São Martinho da Serra station (code MS) is part of the Brazilian SONDA network (<http://sonda.ccst.inpe.br/>),
 installed and administered by the National Institute for Space Research (INPE, Brazil). This network meets the
 quality criteria established by World Meteorological Organization (WMO) and was specifically designed to record
 high-quality meteorological data in different climatic regions of Brazil (Dias da Silva et al., 2014). Cleaning and
 visual inspection at these sites is performed on a daily basis. Based on our experience, the assigned (P95) global
 uncertainty for hourly GHI measurements from these sites (LE, LU, SM) is 3% of the average.

The second group is composed with data from seven sites of Uruguay’s LES solar radiation network, where
 spectrally flat class A or B (according to the new ISO 9060:2018 standard) Kipp & Zonen pyranometers are used
 to measure GHI, among other variables. All these sites are located either at manned meteorological stations or
 agronomic experimental facilities, and the pyranometers are cleaned and inspected at least on a weekly basis.
 These instruments are calibrated at LES at most every two years against the Kipp & Zonen CMP22 secondary
 standard mentioned before. Hourly GHI data from these sites is assigned a typical (P95) global uncertainty of
 5% of the average.

The location of these sites is provided in Table 1 and their geographical distribution is shown in Figure 1.
 The data time-period for each site is provided later in Table 2, jointly with the quality filtering summary.
 Only data sets with a minimum 2-year statistics and complete years (or years and a half) are considered to
 avoid introducing seasonality bias in the data.

2.2. Satellite images

The target area shown in Figure 1 is covered by two geostationary satellites: the GOES-East (operated by
 the National Oceanic and Atmospheric Administration, NOAA) and the MSG (operated by the European
 Organisation

	code	lat (°)	lon (°)	alt (m)
LES facility	LE	-31.28	-57.92	56
São Martinho da Serra	MS	-29.44	-53.82	489
Luján	LU	-34.59	-59.06	30
Canelones (Las Brujas)	LB	-34.67	-56.34	38
Treinta y Tres	TT	-33.28	-54.17	35
Salto	SA	-31.27	-57.89	47
Rocha	RO	-34.49	-54.31	20
Artigas	AR	-30.40	-56.51	136
Colonia (La Estanzuela)	ZU	-34.34	-57.69	70
Tacuarembó	TA	-31.71	-55.83	142

Table 1: Information of the ground measurement stations.

for the exploitation of Meteorological Satellites, EUMETSAT). Due to their positions in the geostationary orbit, they have different pixel sizes and view angles over the area. The GOES-East satellite pixel size is approximately 2 km, as expected over the region for the 1 km nadir spatial resolution of the former GOES12 and GOES13 satellites (Lockheed-Martin, 2019). On the other hand, the MSG satellite has a nadir spatial resolution of 3 km (Schroedter-Homscheidt et al., 2018) and the pixel size over the region is of approximately 7 km¹. The satellites' zenith angles for the target region are approximately 40° and 70° for the GOES-East and MSG, respectively.

Cloud properties and irradiation estimates from satellite images with view angles above 60° are prone to higher errors mainly due to increased pixel size, parallax errors which produce apparent cloud displacement and the failure to fulfill the plane-parallel assumption (Johnson et al., 1994; Schroedter-Homscheidt et al., 2018). CAMS produces regular publicly available validation reports in which its irradiation products are compared to several quality ground sites. Figure 2, based on data from a recent validation report (Lefèvre, 2018), shows the dispersion of the Heliosat-4 estimates (as quantified by rRMSD) vs the satellite zenith angle z of the ground site. A clear threshold is apparent, just below 60°. Sites with $z < 55^\circ$ have average rRMSD of about 11% while those with $z \geq 55^\circ$ have rRMSD of about 25%. So, large viewing angles can affect seriously the accuracy of the irradiation estimates. The CAMS User Manual sets the recommended upper limit for view angle at 60°, while still providing the information for higher view angles (Schroedter-Homscheidt et al., 2018, Sec. 5.2).

Information from both satellites is considered in this work: cloudiness information for the CIMs is derived from GOES-East images, while the Heliosat-4 solar irradiation estimates are based on MSG images. GOES-East satellite images were downloaded from the NOAA CLASS (Comprehensive Large Array-data Stewardship System) website (<https://www.class.noaa.gov/>), where they are publicly available. Information from the MSG satellite is used here out of the recommended zone (satellite zenith angle above 60°) in order to quantify the impact of using such in-

formation for solar resource assessment in the region. The target area location in both satellites fields of view (FOV) is shown in Figure 3.

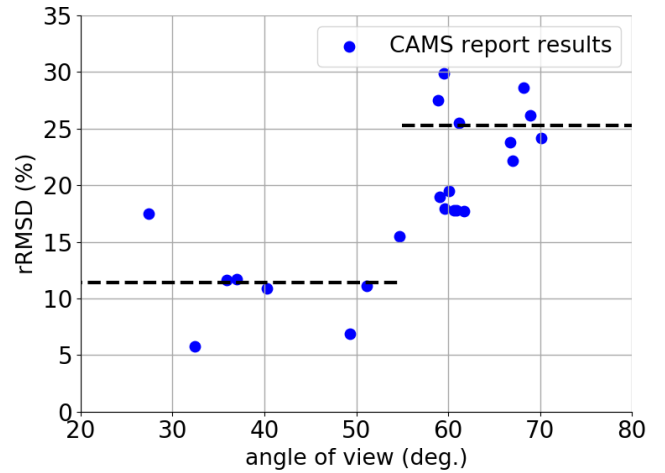


Figure 2: Dispersion (rRMSD) of Heliosat-4 GHI estimates vs view angle z for several sites. Data obtained from (Lefèvre, 2018). The dashed lines indicate average rRMSD for $z < 55^\circ$ and $z \geq 55^\circ$.

2.3. CAMS products

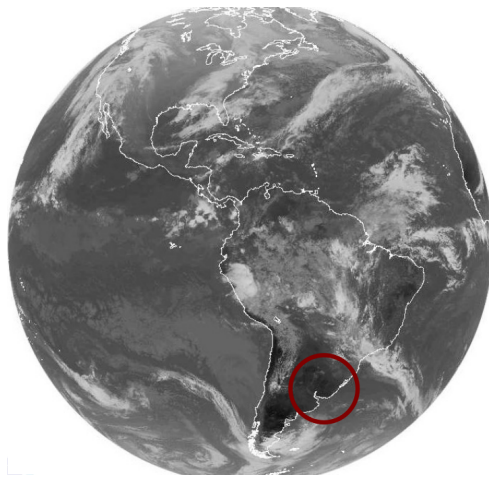
This Subsection briefly describes the CAMS products that are used in this work. The data was retrieved from the CAMS Radiation Service at <http://www.soda-pro.com/>. These estimates are provided with a reliability flag and only the highest reliability estimates were used.

2.3.1. McClear model

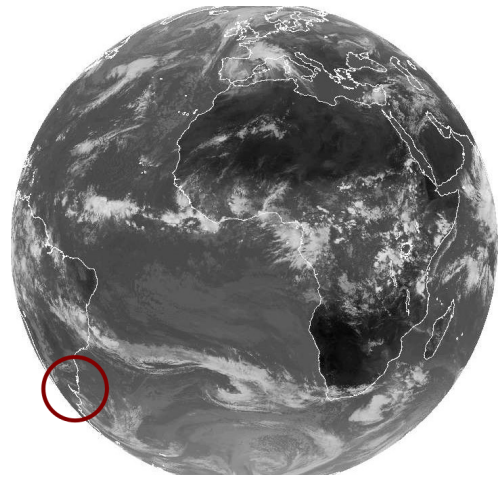
As mentioned in the introduction, the McClear model produces clear-sky GHI estimates based on look-up tables (LUT) of the libRadtran Radiative Transfer Model (RTM) (Lefèvre et al., 2013; Mayer & Kylling, 2005), which in turn uses atmospheric information from satellite retrievals. Being based on LUT, the McClear model can be used operationally (i.e., in real time) since the substantial computational cost of the RTM calculations is avoided. McClear estimates are available at 1-minute intervals with worldwide coverage while the model inputs are typically available every three hours with a spatial resolution between 50-150 km. Using interpolation techniques, the SoDa website provides estimates for any latitude-longitude combination at 1-minute time resolution and above.

In Lefèvre et al. (2013), McClear clear-sky GHI estimates were compared to 1-minute clear-sky measurements from eleven BSRN stations covering different climates in America, Europe, Asia, and Oceania. Mean biases between -1% and $+3\%$ and mean rRMSD in the range 3-5% were obtained (in both cases expressed relative to mean observed irradiance). This model has also been assessed at the 10-minute level against data from seven sites in United Arab Emirates (Eissa et al., 2015b), where the atmosphere is mostly free of clouds but can have high turbidity. The

¹For more information see <http://www.soda-pro.com>



(a) GOES-East satellite FOV. Position: 75°W .



(b) MSG satellite FOV. Position: 0° .

Figure 3: Satellites’ field of view (FOV) identifying in red the area under study. Infrared images are used only for visualization purpose.

319 rMBD was in the range -1% and $+6\%$ and the rRMSD
 320 was between 4-8%. In a similar climate, using 1-minute
 321 measurements from three sites in Israel, McClear had rela-
 322 tive biases between zero and $+4\%$ and a rRMSD of 4%
 323 (Lefevre & Wald, 2016). With the exception of Lefevre
 324 et al. (2013), which considers one site in Brasilia, Brazil
 325 (with different climate and distant more than 2000 km
 326 from the region of interest in this work), there are no other
 327 validations of the McClear model in South America.

328 2.3.2. Heliosat-4 method

329 The Heliosat-4 method (Qu et al., 2017) estimates GHI
 330 and its components under all-sky conditions. It is based
 331 on the McClear model and a second LUT model, the Mc-
 332 Cloud model, also based on RTD calculations. McCloud
 333 estimates the attenuation of Shortwave Solar Irradiation
 334 (SSI) due to cloudiness using a clear-sky index based on
 335 an abacus with four inputs: ground albedo, cloud opti-
 336 cal depth, cloud coverage and cloud type. Cloud infor-
 337 mation is obtained from the APOLLO/SEV methodology,
 338 while ground albedo comes from MODIS. APOLLO/SEV
 339 is an adaptation of NOAA’s APOLLO algorithm (AVHRR
 340 Processing scheme Over cLOUDs, Land and Ocean; Kriebel
 341 et al. (1989, 2003)) for the SEVIRI (Spinning Enhanced
 342 Visible and Infrared Imager) instrument. This procedure
 343 discriminates each pixel in different categories of cloud cov-
 344 erage before deriving its physical properties. The model
 345 considers four categories of cloud type and assigns one
 346 of these types to each covered pixel. The optical depth
 347 is assigned to each cloudy pixel depending on the multi-
 348 spectral APOLLO/SEV procedure that provides this infor-
 349 mation only for fully cloudy pixels. Interpolation tech-
 350 niques are used for pixels in other categories. Based on
 351 this input information, for each abacus node a clear-sky
 352 index is retrieved, and then used to calculate GHI using
 353 a clear-sky libRadtran run over a standard atmosphere.
 354 More details of this sophisticated model can be found in

Qu (2013).

355 The Heliosat-4 method was first validated against mea-
 356 surements from 13 BSRN stations on a 15-minute basis in
 357 Qu et al. (2017). Ten of these sites are located in Europe
 358 (including one in the Canary Islands), while the rest are in
 359 Israel, South Africa and Algeria. This assessment of the
 360 model showed rRMSD values between 15-20% in desert
 361 and mediterranean climates and between 26-43% in rainy
 362 climates with mild winters. The automatic validation re-
 363 ports provided by CAMS at the hourly level ([www.soda-
 364 pro.com/web-services/validation](http://www.soda-pro.com/web-services/validation)) using the period 2014-
 365 2018, include two sites in the target region of this work.
 366 One is an urban site in Buenos Aires (Argentina) and the
 367 other is a coastal site in Florianopolis (Brazil). Neither
 368 of these sites is representative of the region under study.
 369 The former is located in a densely populated urban area
 370 where high atmospheric turbidity is frequent. The latter
 371 is located in an Atlantic coast island with a climate dom-
 372 inated by the ocean and is more than 1000 km away from
 373 the closest site used in this work. Performance in these
 374 two sites shows rMBDs of -5% and 0% and rRMSDs of
 375 25% and 28%, respectively. Both sites have relatively high
 376 viewing angles outside the recommended range, 73° for
 377 Buenos Aires and 62° for Florianopolis. In Toravere, Es-
 378 tonia, a region with similar satellite view angle (70°), the
 379 evaluation shows a positive bias of $+3\%$ and an rRMSD
 380 of 28%. On the other hand, in Carpentras, France, with
 381 zenith angle 51° the rMBD $+2\%$ and the rRMSD is only
 382 15%. Another case is Tamanrasset, Algeria, a desertsic lo-
 383 cation near the satellite nadir (zenith angle 27°), where the
 384 performance is -5% rMBD and 15% rRMSD. The rRMSD
 385 for these and other sites are included in the CAMS quar-
 386 terly validation report and are plotted vs satellite zenith
 387 angle in Figure 2 to display the increase in rRMSD for
 388 sites beyond $z = 60^{\circ}$.
 389

390 3. Methodology

391 This Section is organized as follows. Subsection 3.1
392 presents the quality assessment and clear-sky selection pro-
393 cedures applied to the data sets. Subsection 3.2 describes
394 the locally implemented models, ESRA model and both
395 CIMs, including the cloud index calculation and the spatial
396 smoothing applied to the satellite information to reduce
397 the uncertainty of hourly estimates. This spatial smooth-
398 ing procedure is also applied to the CAMS Heliosat-4 prod-
399 uct, for a fair comparison, as explained at the end of Sub-
400 section 3.2.4. Finally, Subsection 3.3 describes the local
401 adaptation procedures used for the different models and
402 estimates. This includes the adjustment of locally imple-
403 mented models and the site adaptation of CAMS products,
404 for fair comparison.

405 3.1. Pre-processing of data

406 Hourly horizontal irradiation was calculated for each
407 site from the original one-minute ground measurements.
408 Hours with more than 10-minute gaps were discarded. The
409 following quality-control filters were applied to the day-
410 light hourly GHI measurements:

- 411 (i) Minimum solar elevation: $\alpha_s > 7^\circ$; in order to avoid
412 data affected by large cosine errors.
- 413 (ii) Maximum irradiation: $\text{GHI} < \text{GHI}_{\text{csk}}^*$; the ESRA
414 clear-sky model (Subsection 3.2.1) was used with a
415 low Linke turbidity factor $T_L = 2$ to compute $\text{GHI}_{\text{csk}}^*$
416 as an upper bound for GHI. This value of T_L is suf-
417 ficiently low (see the average T_L cycle shown in Fig-
418 ure 5) for this purpose and it has been previously
419 used in this region to generate an upper bound for
420 hourly GHI (Abal et al., 2017).
- 421 (iii) Modified clearness index bounds: $0 < k'_T < 0.85$;
422 the modified clearness index, k'_T , is defined in Perez
423 et al. (1990).
- 424 (iv) Coincident pairs of GHI and GOES-East satellite in-
425 formation. The samples discarded at this stage are
426 mainly determined by the GOES-East satellite avail-
427 ability, as discussed below.
- 428 (v) Coincident pairs of GHI and Heliosat-4 estimates.
429 As mentioned, only Heliosat-4 estimates flagged with
430 the highest reliability are considered.

431 GOES-East image availability for South America was
432 irregular before the year 2018, when the new GOES-R
433 started operations at the GOES-East position. For the
434 period 1997-2017, these images are normally available at
435 a rate of two per hour. However, hourly or tri-hourly gaps
436 result for South America when the GOES-East satellite
437 was placed under Rapid Scan Operation mode for spe-
438 cific areas. The hourly satellite information was obtained
439 by linear interpolation of the satellite time series under

the restriction of not interpolating across gaps larger than
three hours.

440 The filtering results for each site are summarized in
441 Table 2, where filters (ii) and (iii) are grouped in a sin-
442 gle column for brevity. It describes the filters sequentially
443 applied to the initial data set (daylight values), so that
444 each discard percentage refers to the previous column and
445 the number of records that passed all previous filters is
446 informed at each stage. The last two columns indicate the
447 final hourly records for all-sky and clear-sky conditions,
448 respectively. After this filtering procedure, a set of 160298
449 hourly GHI records are available for the all-sky model as-
450 sessment.
451

452 The selection of the clear-sky subset was based on the
453 procedure proposed by Remund et al. (2003). This algo-
454 rithm is based on five consecutive filters applied to hourly
455 data, but the main criterion is to impose a lower thresh-
456 old of 0.7 on the modified clearness index, k'_t (Perez et al.,
457 1990). We added an extra filter to this procedure imposing
458 a bound on daily variability: if the standard deviation of
459 the k'_t series within a day was over 0.05, the whole day was
460 discarded. The threshold of 0.05 was heuristically deter-
461 mined to ensure that only clear-sky records were selected,
462 since any contamination by partly cloudy samples would
463 artificially affect the clear-sky models performance assess-
464 ment. The amount of clear-sky hours selected for each site
465 is indicated in the last column of Table 2. Considering all
466 sites, 34050 hourly clear-sky samples were selected.
467

468 3.2. Locally implemented models

469 3.2.1. ESRA model

470 The ESRA clear-sky model was developed in the frame-
471 work of the European Solar Radiation Atlas (Rigollier et al.,
472 2000) and used with Meteosat images as part of the Heliosat-2
473 method for SSI modeling (Rigollier et al., 2004). It esti-
474 mates direct normal irradiance (DNI) and diffuse horizon-
475 tal irradiance (DHI) under clear-sky conditions. The GHI
476 estimate is obtained from $\text{GHI} = \text{DNI} \times \cos \theta_z + \text{DHI}$. The
477 single input in this model is the Linke Turbidity factor,
478 T_L , for air mass 2. It is usually interpreted as the num-
479 ber of clean, dry atmospheres (i.e. with no clouds, water
480 vapor or aerosols) which would produce the same attenua-
481 tion effect on GHI as the real cloudless atmosphere. Thus,
482 T_L includes in one effective parameter the information on
483 water vapor density and aerosol contents of the real atmo-
484 sphere (Linke, 1922; Ineichen & Perez, 2002). Given its
485 simplicity and the fact that it may provide accurate esti-
486 mations if the T_L values are locally obtained with sufficient
487 time resolution (Gueymard, 2012), the ESRA model is a
488 frequent choice to model clear-sky irradiation. Here, the
489 T_L cycles have been derived on an monthly basis from the
490 GHI measurements, as discussed in Subsection 3.3.

491 ESRA model performance has been analyzed in several
492 studies (Gueymard, 2012; Engerer & Mills, 2015; Ineichen,
493 2016; Antonanzas-Torres et al., 2019; Sun et al., 2019) that
494 consider different climates around the world. This model

site	period	daylight	(i) solar altitude		(ii) & (iii) bounds		(iv) GOES images		(v) CAMS reliability		clear-sky
		hours	disc.	hours	disc.	hours	disc. (%)	hours	disc.	hours	hours
LE	01/15–12/17	12432	9.8%	11216	1.9%	11004	0.4%	10964	6.0%	10307	2783
MS	01/10–12/16	30199	9.0%	27493	2.3%	26854	1.8%	26367	6.9%	24553	5257
LU	01/10–12/13	17025	10.7%	15195	3.0%	14741	3.1%	14283	9.5%	12925	3282
LB	06/10–12/17	28826	10.3%	25843	1.8%	25383	2.0%	24882	6.6%	23232	4931
TT	06/10–05/16	24705	9.8%	22294	3.0%	21629	2.2%	21153	7.1%	19652	2790
SA	06/10–12/14	19083	9.8%	17216	1.7%	16930	2.9%	16446	8.1%	15122	4412
RO	06/11–12/17	27120	9.6%	24514	2.8%	23831	2.0%	23351	6.0%	21940	3329
AR	01/12–12/17	21955	9.0%	19970	2.3%	19509	1.1%	19290	6.3%	18071	4239
ZU	01/16–12/17	8686	10.6%	7766	1.6%	7645	0.5%	7605	5.2%	7206	1514
TA	01/16–12/17	8659	9.1%	7871	2.2%	7697	0.5%	7658	4.8%	7290	1513
total		198690	9.7%	179378	2.3%	175223	1.8%	171999	6.8%	160298	34050

Table 2: Quality check and data set description for each ground measurement site. The % discarded and the number of hours that pass each filter are informed. The last two columns indicate the all-sky and clear-sky hours used for model assessment.

usually is well ranked among other simple proposals and its uncertainty mostly depends on the quality of the T_L input data being used. Its validation in the Pampa Húmeda region has been scarce and comprises only one preliminary local study using five measurement sites in Uruguay (Laguarda & Abal, 2017) where a rMBD of -0.5% and a rRMSD of 4.5% were found for clear-sky hourly GHI estimation. Validations for similar climates (Cfa and Cfb in the Köppen-Geiger classification) are as follows. In Gueymard (2012) a rMBD of $+4.3\%$ and rRMSD of 4.9% was found for the ARM-SGP site (Oklahoma, USA) using 1-minute GHI measurements. In Engerer & Mills (2015) the ESRA model was evaluated at 14 sites using 1-minute GHI data from the Australian Bureau of Meteorology, four of which are in the relevant climate zones. For these, rMBDs between $+2\%$ and $+9\%$, and rRMSDs between 3.7% and 8.0% were found. An exhaustive revision of 38 validation studies of clear-sky models performance, most of them including ESRA, can be found in Ruiz-Arias & Gueymard (2018). Overall, there is a considerable spread in the performance of the ESRA model, depending on climate and implementation details, and simplicity is one of its key features.

3.2.2. Cloud Index Methods (CIM)

As mentioned before, this family of SSI models has the common structure of a clear-sky model with a modulating factor that takes into account the effect of clouds. The clear-sky index, $\text{GHI}/\text{GHI}_{\text{csk}}$, can be modeled by a cloud attenuation factor, $F(C)$, which depends on the satellite-derived cloud index C defined in the following in Eq. (3). In this work, we use a simple linear function,

$$F(C) = a + b \times (1 - C), \quad (1)$$

where a and b are locally adjusted for each site. Then, GHI is computed from:

$$\text{GHI} = \text{GHI}_{\text{csk}} \times F(C). \quad (2)$$

In this work, Eqs. (1) and (2) are implemented using the two clear-sky models discussed before (ESRA and McClear) and the resulting CIMs are referred in what follows

as CIM-ESRA and CIM-McClear, respectively. We emphasize that the a and b parameters in Eq. (1) are site and model-specific. The coefficients for CIMs in the region are part of the results of this work and are discussed in Subsection 3.3 on local adaptation (Table 3).

3.2.3. Cloud index calculation

The satellite-derived cloud index (Cano et al., 1986), C , is a dimensionless parameter in $[0, 1]$ that quantifies the amount of cloudiness. It is obtained from the Earth albedo (or planetary reflectance), ρ_p , by normalization with extreme values ρ_{\min} and ρ_{\max} associated with clear and overcast skies respectively,

$$C = \frac{\rho_p - \rho_{\min}}{\rho_{\max} - \rho_{\min}} \quad \text{for } \rho_{\min} < \rho_p < \rho_{\max}. \quad (3)$$

Additionally, the constraints $C = 1$ for $\rho_p > \rho_{\max}$ and $C = 0$ for $\rho_p < \rho_{\min}$ are imposed. The parametrization proposed in Tarpley (1979) is used in this work to estimate the intra-day and seasonal variation of the background albedo. This parametrization models the background reflectance factor, $F_{Ro} = \rho_{po}/\cos\theta_z$, and needs to be adjusted for each pixel in the image using satellite clear-sky samples. These clear-sky samples are automatically selected from the pixel's satellite time-series by a robust iterative procedure described in Alonso-Suárez et al. (2012). This adjustment procedure can be updated on real time taking the past pixel samples, but for the sake of this work it was done only one time using the 2010–2017 satellite period. After the coefficients for each pixel (or site) are adjusted, whether on real-time or offline, the parametrization can be used to estimate the F_{Ro} time-series at hourly intervals. Then, the ρ_{\min} time-series is calculated by setting $\rho_{\min} = \rho_{po} = F_{Ro}/\cos\theta_z$. A proper background albedo characterization is an important step in order to obtain useful cloud index information. For ρ_{\max} , a fixed constant value of 0.80 is chosen, since this value has been found to optimize the performance in the region of satellite-based models for GHI (Laguarda et al., 2018).

3.2.4. Spatial smoothing

In order to use satellite information at an hourly basis, the reflectance ρ_p obtained from GOES-East images has been spatially averaged in a $10 \text{ min} \times 10 \text{ min}$ latitude-longitude cell centered at the site of interest. For this target region, this corresponds to cells of approximately $16 \text{ km} \times 18 \text{ km}$. This is equivalent to an ergodic hypothesis, where the spatial average of an instantaneous image within a cell is representative of the average conditions at the center of the cell within the hour. The cell size has been optimized to minimize the uncertainty of hourly satellite models. The rRMSD trend as a function of the cell size is shown in Figure 4, based on the CIM-McClear model. This curve is essentially the same as a similar one reported in Alonso-Suárez (2017), but using different ground stations, data time-span and satellite model. Hence, this curve can be considered characteristic for locally-adapted satellite based models in the target region. Inspection of Figure 4 shows that the minimum is shallow and essentially the same performance is obtained between 8 min and 12 min latitude-longitude spacing, which approximately correspond to cells with $12 \text{ km} \times 15 \text{ km}$ and $18 \text{ km} \times 22 \text{ km}$ area, respectively.

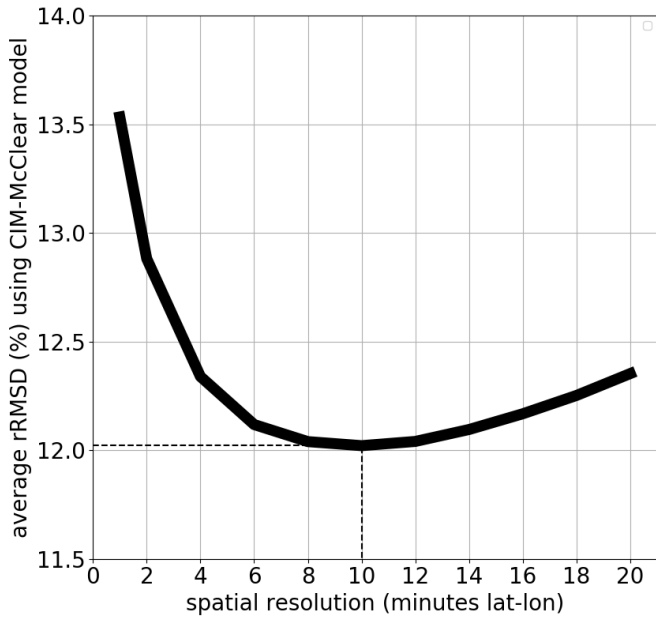


Figure 4: Local rRMSD curve as a function of the spatial smoothing.

The Heliosat-4 estimates are generated for single pixels without any spatial smoothing (pixel size is around 6-7 km, as discussed in Subsection 2.2). For fair comparison, the Heliosat-4 estimates were downloaded for each site in a grid of 3×3 pixels surrounding each location, accounting for a similar spatial averaging ($\approx 21 \text{ km} \times 21 \text{ km}$). The average of the nine time-series has been used as the model's estimates for each site.

3.3. Local adaptation

The aim of local adaptation is to reduce bias and, more generally, to improve model performance in a given homogeneous geographical area. It can be achieved either by locally adjusting the model's parameters or by site-adapting their estimates. Since the CIMs have locally adjusted parameters, for a fair comparison, the estimates available from CAMS for the McClear and Heliosat-4 models must be site-adapted.

A widely used site adaptation technique consists of a linear regression correction between the hourly model estimates and the ground data (Polo et al., 2016). This strategy is used in this work to site-adapt the estimates from both CAMS models. In Section 4 the results for McClear model and Heliosat-4 method are provided with and without site-adaptation, so the performance gain of the adaptation procedure can be observed.

CIMs, based on the site-adapted McClear or ESRA estimates, are locally adapted by adjusting their two parameters to ground data (a and b , see Eq. (1)) using a standard cross validation technique, where half of the data is randomly selected to train the model and the other half is used to evaluate its performance. The procedure is repeated 1000 times to ensure repeatability, and the ensemble average uncertainty and adjusted parameters are reported. In this context, it is worth pointing out that the cloud index calculation includes also an implicit local adaptation at each pixel, as the ground albedo (image background brightness) has been locally adjusted by modelling the F_{Ro} (and ρ_{min}) time series, as described in Subsection 2.2. The locally adjusted parameters for the CIMs are shown in Table 3 for each site. Both CIMs have similar values for a and b and there is a good agreement across sites. This is considered a sanity check for the proposals and the adjustment, as the region is mostly uniform in geography and climate. Thus, the average set of parameters can be used for the region without significantly affecting performance.

In the case of the ESRA clear-sky model, the local adjustment is made through the T_L values used as input. Yearly cycles of average T_L values are estimated for each site using its GHI data and the ESRA GHI clear-sky parametrization. These values were obtained from clear-sky samples of ground measurements by minimizing the statistical deviation between the model and the ground truth, as detailed in Laguarda & Abal (2016). The resulting average T_L cycles have a small spatial variation, as shown in Figure 5, where the spatially averaged yearly cycles for T_L are shown for two broad regions which correspond approximately to the areas of Figure 1 separated by latitude 33°S (North and South). Values for T_L are between 2.8 and 4.2, with higher values in the summer and lower values in winter.

This method captures seasonal effects and models the average trends in the local atmospheric turbidity and water vapor. However, it does not attempt to model its daily or hourly variability. In this way, the issue of using dif-

site	CIM-McClea		CIM-ESRA	
	a	b	a	b
LE	0.073	0.92	0.040	0.94
MS	0.076	0.91	0.053	0.92
LU	0.062	0.92	0.041	0.93
LB	0.068	0.91	0.044	0.93
TT	0.072	0.90	0.052	0.91
SA	0.055	0.93	0.031	0.94
RO	0.075	0.90	0.054	0.91
AR	0.076	0.91	0.048	0.92
ZU	0.070	0.91	0.045	0.92
TA	0.074	0.91	0.045	0.93
mean	0.070	0.92	0.045	0.93
σ	9.7 %	1.0 %	15.6 %	1.2 %

Table 3: Locally adjusted parameters of cloud index models (Eq. (1)). The last two rows show the weighted average and the standard deviation (P67) as a %.

648 ferent T_L formulae based on different quality satellite re-
649 trievals or atmospheric inputs is avoided.

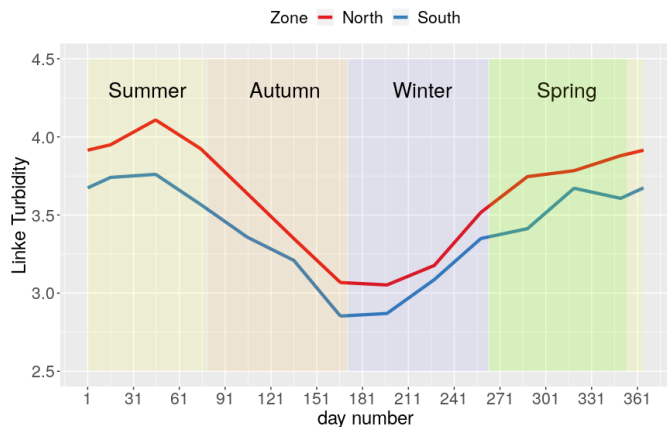


Figure 5: Daily cycles for T_L (no unit) obtained from clear-sky data in Laguarda & Abal (2016). The northern and southern zones are separated approximately by the 33°S latitude parallel (see Figure 1).

650 4. Results

651 4.1. Performance metrics

652 The performance assessment is done using three com-
653 mon indicators: the mean bias deviation (MBD), the root
654 mean square deviation (RMSD) and the Kolmogorov-Smirnov
655 integral (KSI). The first two measure the average bias and
656 the average dispersion of the residuals, respectively. These
657 are expressed in relative terms as a percentage of the mea-
658 surement average (rMBD and rRMSD, respectively). The
659 KSI is a statistical similarity index based on the distance
660 between the probability distributions of the measurements
661 and the estimates (Massey Jr., 1951; Espinar et al., 2009).

662 A useful discussion and examples of use of these indicators
663 can be found in Gueymard (2014).

664 Whenever average metrics over all sites are reported,
665 the P95 uncertainty assigned to each ground measure-
666 ment’s data set is used to weight the averages, so that
667 higher quality data will have more impact on the indica-
668 tors. The weight for each site is calculated as $w_i = c/u_i^2$
669 with u_i the assigned relative uncertainty for measurements
670 from site i . The set of weights is scaled by c to add up
671 to unity, $c \times \sum_i 1/u_i^2 = 1$. This standard weighting pro-
672 cedure has been previously used in a similar context with
673 good results (Abal et al., 2017).

674 4.2. Clear-sky models (McClea, ESRA)

675 The performance assessment for clear-sky models un-
676 der cloudless conditions is shown for each site in Table 4,
677 where the last column shows the (weighted) average indi-
678 cators over all sites and the last row shows the ground
679 measurement averages for the clear-sky hours under com-
680 parison. This information is provided to enable the reader
681 to compute the absolute indicators, if needed.

682 Both models, whether locally-adapted or not, perform
683 well and within the expected ranges at all sites, and the
684 rRMSD indicators are similar to the P95 uncertainty as-
685 signed to the high quality ground data sets. The McClea
686 estimates as provided by the CAMS platform (without site
687 adaptation) show a small positive average bias of +1.4%
688 with small but consistent overestimation at all sites. The
689 rRMSD values range from 2.7% to 4.3% with an average
690 of 3.2% while the average KSI is 8.9 Wh/m². This non
691 site-adapted model’s performance is similar to that of the
692 ESRA model with locally adjusted T_L values, which has
693 an average rRMSD of 3.5%. However, the site variabil-
694 ity is lower for the ESRA model, ranging from 3.2% to
695 3.8%. This model is essentially unbiased (its average bias
696 is -0.1% and it is within $\pm 0.3\%$ across all sites). This
697 results in a lower KSI metric than the original McClea
698 estimates, with a site-averaged KSI of 5.0 Wh/m².

699 The site-adapted McClea model provides the best per-
700 formance: it is unbiased and consistently has the lowest
701 rRMSD and KSI across all sites. The site-average rRMSD
702 of the locally adapted McClea is 2.8%, ranging from 3.7%
703 at the oceanic RO site and as low as 2.5% at the high
704 quality LE site. Similarly, the average KSI is 1.9 Wh/m²,
705 showing a superior performance also from a statistical sim-
706 ilarity point of view. As mentioned, this model takes into
707 account the atmospheric short-term variability, while the
708 less sophisticated ESRA model only takes into account sea-
709 sonal trends in atmospheric turbidity.

710 These indicators (low or negligible bias deviation and
711 rRMSD in the range 3-4%) are not surprising from lo-
712 cally adjusted clear-sky models (Gueymard, 2012). Lower
713 indicators (around 2%) have been reported for detailed
714 models with high quality atmospheric information, such as
715 the REST2 clear-sky model at particular locations (Guey-
716 mard, 2008). Taking into account the uncertainties of our
717 hourly ground data set, the performance assessment of the

718 clear-sky models cannot be made to those limits. A shorter
719 1-minute time scale and pyrheliometer data (accurate to
720 1%) would be required to further explore the performance
721 limits of the McClear or other models, such as REST2, in
722 this region. However, this is out of the scope of this work.

723 4.2.1. Clear-sky models under all-sky conditions

724 In several applications, such as CIMs or short-term
725 forecasting, the output of a clear-sky model is used under
726 non clear-sky conditions. Thus, it is relevant to investigate
727 the output characteristics of clear-sky models under non
728 clear-sky conditions.

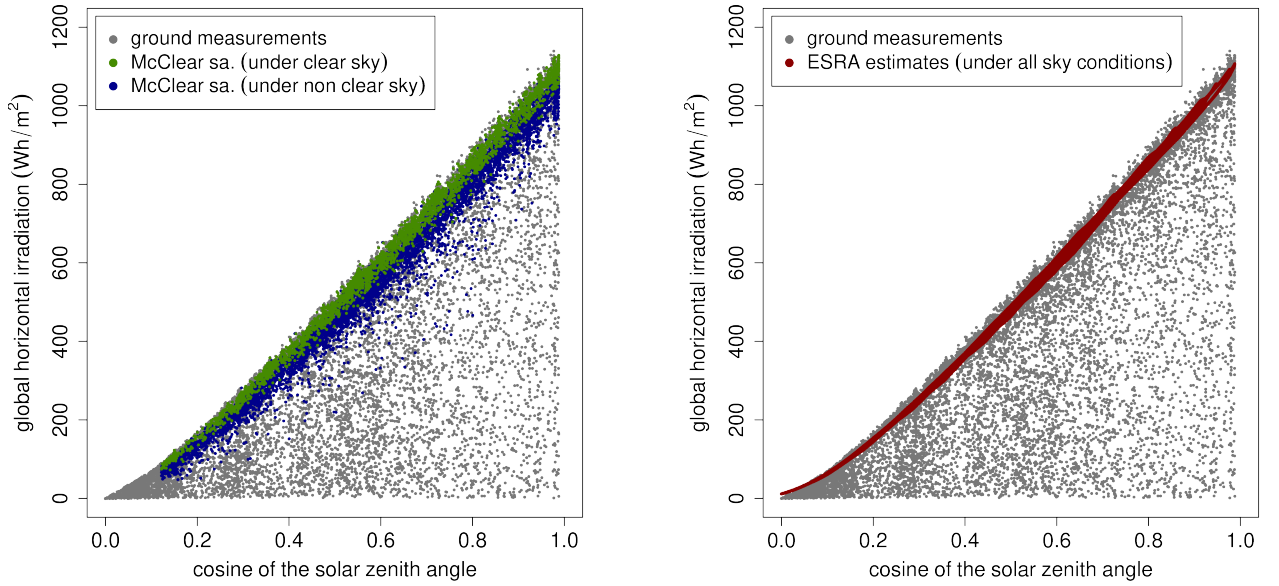
729 **Figure 6** reveals a relevant difference between the two
730 clear-sky models estimates under non clear-sky conditions.

731 Both panels show the hourly clear-sky estimates from each
732 model as a function of the cosine of the solar zenith angle
733 for all-sky conditions. The ground measurements (all-sky)
734 are shown in the background in grey. As shown in **Fig-**
735 **ure 6a**, the estimates from the McClear model are affected
736 by the actual sky condition, being lower when there are
737 clouds in the real atmosphere (blue dots) than under real
738 clear-sky conditions (green dots). This behavior is not
739 observed in the ESRA model (**Figure 6b**) where the esti-
740 mates under clear-sky and all-sky condition show the same
741 characteristics.

742 In the presence of clouds, particularly under heavy
743 overcast conditions, there is more water vapour in the at-
744 mosphere. Since the McClear model takes into account the

model	metric	LE	MS	LU	LB	TT	SA	RO	AR	ZU	TA	all sites
McCclear (original)	rMBD (%)	1.1	1.8	0.7	1.7	1.8	1.3	2.1	2.7	2.1	2.8	1.4
	rRMSD (%)	2.7	3.4	3.0	3.6	3.8	3.1	4.3	4.1	3.6	4.0	3.2
	KSI (Wh/m ²)	6.8	11.5	4.6	10.3	10.5	7.9	13.3	17.3	13.3	17.6	8.9
McCclear (site adapted)	rMBD (%)	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
	rRMSD (%)	2.5	2.8	2.9	3.2	3.3	2.8	3.7	3.0	2.9	2.9	2.8
	KSI (Wh/m ²)	1.5	1.7	2.1	1.9	2.7	2.5	2.0	1.8	3.6	2.4	1.9
ESRA (adjusted T_L)	rMBD (%)	-0.1	-0.2	0.1	-0.1	-0.1	-0.1	-0.1	-0.1	0.2	-0.3	-0.1
	rRMSD (%)	3.2	3.6	3.5	3.4	3.7	3.4	3.8	3.7	3.6	3.4	3.5
	KSI (Wh/m ²)	4.2	5.6	4.9	4.6	5.1	5.1	5.2	5.6	5.7	5.4	5.0
measurement average (Wh/m²)		641	632	621	610	595	621	626	629	622	640	629

Table 4: Performance of the McClear original model, the site-adapted McClear model and the locally adjusted ESRA model under clear-sky conditions. The last row is the measurement average of the clear-sky hours. The last column holds the weighted metrics average over sites.



(a) Site-adapted McClear estimates under clear-sky (green) and non clear-sky (blue) conditions.
(b) Locally-adapted ESRA estimates under all-sky conditions (clear-sky and non clear-sky conditions).

Figure 6: Clear-sky estimates under different real sky conditions for McClear and ESRA models vs the cosine of the solar zenith angle. In the background, the measured ground data is in grey. For colors please refer to the online version of the manuscript.

745 actual water vapour content, it provides lower clear-sky estimates
746 under cloudy sky condition than under clear-sky conditions. On the other hand, the ESRA model as im-
747 plemented here is based on T_L monthly averages, so it is
748 insensitive to short term variations in the atmosphere, in
749 particular to the presence of clouds. This difference in
750 the behaviour observed for these two models may appear
751 when any model with real-time atmospheric water vapor
752 information is compared to a model which uses average
753 information for a given region. This issue should be taken
754 into account in applications using clear-sky estimates under
755 all-sky conditions.

756 To quantify this difference, a comparison is made be-
757 tween the clear-sky estimates from both models, stratified
758 according to the actual sky condition. The average deviations (rMD)
759 between both models are expressed according
760 to $\Delta = \text{GHI}_{\text{csk}}^{\text{mcclear}} - \text{GHI}_{\text{csk}}^{\text{esra}}$, as the ESRA model estimates
761 are not affected by the presence of clouds and thus
762 it is more similar to a fixed upper limit for hourly all-sky
763 conditions (see Figure 6b). The comparison between both
764 models, shown in Table 5, has a negligible mean deviation
765 under clear sky conditions (+0.1%, on average). How-
766 ever, in the presence of cloudiness, McClear estimates are
767 systematically lower than ESRA estimates, with an average
768 difference of -5.2% . Furthermore, negative deviations
769 ranging between -3.9% and -6.9% are observed consistently
770 across sites. It is important to emphasize that this
771 analysis does not imply ranking one model over the other
772 in terms of accuracy, but rather to highlight and quantify
773 their different behaviour under cloudy conditions.

774 4.3. All-sky models (Heliosat-4, CIM)

775 Results for the all-sky models are presented in this Sec-
776 tion. The evaluation includes four models: the original
777 Heliosat-4 estimates, the site-adapted Heliosat-4 estimates
778 and the locally adjusted CIM-ESRA and CIM-McCclear
779 models. The inclusion of the original Heliosat-4 estimates
780

781 allows an evaluation of the impact of site-adaptation for
782 this model in the region. The performance evaluation for
783 the four all-sky models is presented in Table 6.

784 The original Heliosat-4 estimates have a low overall
785 bias deviation of -0.8% , but with a variation across sites
786 within $\pm 1.9\%$. The average rRMSD metric is of 17.9% ,
787 which is to be expected for a non locally adapted and space
788 averaged satellite based model. High dispersion across
789 sites is also seen in the KSI metric, ranging from about
790 10.1 to 23.4 Wh/m^2 with an average value of 17.7 Wh/m^2 .
791 In the unbiased site-adapted version, the rRMSD metric
792 is reduced only slightly to 16.8% while the KSI becomes
793 10.2 Wh/m^2 (average values). The KSI is a more sensitive
794 indicator than rRMSD and it is more reduced by the site
795 adaptation procedure.

796 The locally adjusted models (CIM-ESRA and CIM-
797 McCclear) have small negative biases across all sites (neg-
798 ative and less than 1.4% in absolute magnitude) with a
799 weighted average of -1.1% in both cases. This is a small
800 but consistent underestimation observed for these models
801 across all sites. The rRMSD metric ranges from 11.2%
802 to 14.1% for the CIM-ESRA and from 10.6% to 13.8%
803 for the CIM-McCclear. The average rRMSD is 12.5% for
804 CIM-ESRA and 12.1% for the CIM-McCclear model. This
805 represents a significant improvement with respect to the
806 site-adapted Heliosat-4 estimates, whose rRMSD is in the
807 15.4 - 19.6% range, with an average of 16.8% . Similarly,
808 the average KSI is 7.0 Wh/m^2 for the CIM-ESRA and
809 8.4 Wh/m^2 for the CIM-McCclear which also represents
810 a reduction with respect to the site-adapted Heliosat-4
811 (10.2 Wh/m^2 on average). Overall, the CIMs represent
812 a significant improvement in the accuracy of GHI estima-
813 tion for this region, with respect to the Heliosat-4 method
814 whether it is site-adapted or not. As the Heliosat-4 is a
815 sophisticated method that considers many atmospheric in-
816 puts and phenomena, usually ranked among the best per-
817 forming models, these important difference of more than

site	average of ESRA estimates (Wh/m^2)	rMD (%)		rRMSD (%)		KSI (Wh/m^2)	
		csk	non csk	csk	non csk	csk	non csk
LE	599	+0.1	-6.9	2.8	9.8	3.9	41.4
MS	608	+0.2	-4.9	2.6	7.6	5.0	29.8
LU	579	-0.1	-4.0	2.6	5.7	4.2	23.0
LB	566	+0.1	-4.2	2.8	6.5	4.6	23.8
TT	582	+0.1	-3.9	2.3	5.9	4.2	22.8
SA	589	+0.1	-5.0	2.8	7.2	5.1	29.7
RO	576	+0.1	-3.9	2.6	6.0	4.1	22.4
AR	592	+0.1	-5.8	2.8	8.4	4.9	34.4
ZU	579	-0.1	-4.9	2.8	7.7	5.2	28.4
TA	586	+0.3	-5.7	2.7	8.4	4.6	33.1
average	592	+0.1	-5.2	2.7	7.6	4.4	30.6

Table 5: Comparison between clear-sky models estimates: site-adapted McCclear vs ESRA, under clear-sky conditions (csk) and non clear-sky conditions (non cks). The higher ESRA average is used for normalization, without implying greater accuracy. The last row shows the weighted average across sites.

818 4% of rRMSD metric in comparison with simple CIMs is
819 mostly attributable to the use of the different satellite in-
820 formation, particularly, the different satellite FOV for the
821 region. Local adaptation and spatial smoothing have been
822 done to both type of models, so the difference is not at-
823 tributable to these features. Heliosat-4 estimates are used
824 here outside the recommended area and its performance is
825 affected by the large viewing angle of the MSG satellite in
826 the region. Hence, the difference found is a quantification
827 of the increase in uncertainty when using CAMS all-sky
828 GHI estimates out of the recommended area.

829 The CIM-McClear model has a slightly better perfor-
830 mance than the CIM-ESRA in terms of rRMSD (12.1%
831 compared to 12.5%). However, the opposite behavior is
832 observed in statistical similarity, with an average KSI be-

833 ing 7.0 Wh/m^2 for the CIM-ESRA and 8.4 Wh/m^2 for the
834 CIM-McClear. Therefore, it can be concluded that both
835 CIMs have a remarkable good performance in this region
836 and both can be used with low uncertainty for solar re-
837 source assessment. However, if these metrics are priori-
838 zed in the presented order (MBD, RMSD and KSI), the
839 CIM-McClear model performs slightly better, as it has a
840 similar bias but a reduced RMSD in comparison with CIM-
841 ESRA.

842 The scatter plots in Figure 7 compare the three locally
843 adapted all-sky models with the ground measurements for
844 all sites data. The smaller dispersion of the two CIMs
845 can be seen with the naked eye. At higher irradiances
846 ($\text{GHI} > 800 \text{ Wh/m}^2$) an overestimation is apparent in
847 the site-adapted Heliosat-4 estimates. For the same con-

model	metric	LE	MS	LU	LB	TT	SA	RO	AR	ZU	TA	all sites
Heliosat-4 (original)	rMBD (%)	-1.0	0.6	-3.1	-1.5	0.8	-2.6	2.6	0.2	0.6	1.9	-0.8
	rRMSD (%)	17.5	19.4	16.6	17.3	17.8	18.0	20.4	16.8	17.7	17.7	17.9
	KSI (Wh/m^2)	22.0	12.8	21.3	16.0	10.1	23.4	11.2	13.6	13.9	12.5	17.7
Heliosat-4 (site adapted)	rMBD (%)	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
	rRMSD (%)	16.0	18.6	15.4	16.4	17.3	16.5	19.6	16.0	16.9	17.0	16.8
	KSI (Wh/m^2)	9.5	11.6	10.3	9.5	10.1	9.0	10.6	9.7	8.8	9.0	10.2
CIM-ESRA (adjusted T_L)	rMBD (%)	-0.9	-0.9	-1.2	-1.1	-1.2	-1.1	-1.3	-1.1	-1.1	-1.3	-1.0
	rRMSD (%)	11.2	14.1	12.3	12.2	13.1	11.8	13.6	11.4	11.9	12.0	12.5
	KSI (Wh/m^2)	5.1	6.3	9.3	7.8	9.0	6.5	8.9	6.2	6.6	6.7	7.0
CIM-McClear (site adapted)	rMBD (%)	-0.6	-1.2	-1.3	-1.1	-1.3	-1.1	-1.4	-1.0	-1.2	-1.0	-1.1
	rRMSD (%)	10.6	13.8	12.1	12.0	12.9	11.3	13.6	11.0	11.5	11.6	12.1
	KSI (Wh/m^2)	4.9	9.2	10.8	8.7	10.7	7.9	10.7	7.2	8.5	7.0	8.4
measurement average (Wh/m^2)		463	446	465	438	437	469	427	458	440	438	448

Table 6: Performance metrics for the all-sky models against ground measurements: original Heliosat-4, site-adapted Heliosat-4, CIM based on ESRA and CIM based on site-adapted McClear. The last column shows the weighted average across sites.

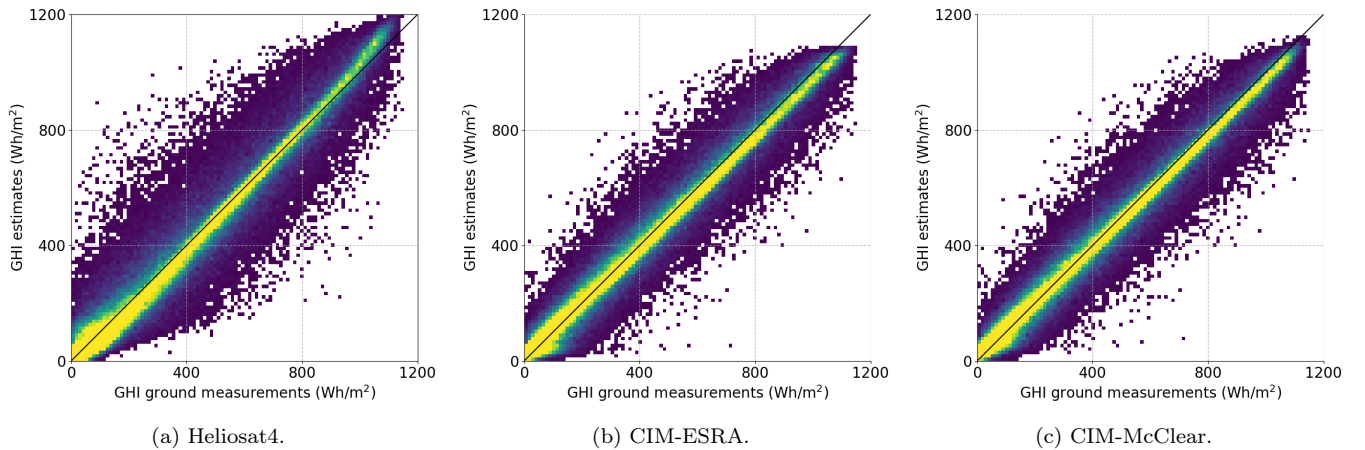


Figure 7: Estimates vs ground measurements for the site adapted Heliosat-4 (left), and the cloud index models based on ESRA (center) and the site adapted McClear (right) for all sites. The colour scale indicate the concentration of the samples in the scatter plot.

848 ditions, a small underestimation bias is observed in both
 849 CIMs. Also, a different behaviour is observed at low irra-
 850 diation between the site-adapted and the locally adjusted
 851 models, with a small overestimation tendency for the former.
 852 For high irradiation values, the CIM-ESRA model
 853 present limitations, as it can be seen in the upper right
 854 side of [Figure 7b](#). This is a consequence of using the ESRA
 855 clear-sky model as basis with an average yearly T_L cycle.

856 To throw some light on the specific shortcomings of
 857 each model under the different conditions, their perfor-
 858 mance indicators are discriminated by clearness index, $k_T =$
 859 GHI/GHI_0 (where GHI_0 is the extraterrestrial solar ir-
 860 radiation at a horizontal plane at the top of the atmo-
 861 sphere) and by the cosine of the solar zenith angle, $\cos\theta_z$
 862 ([Iqbal, 1983](#)). The resulting diagrams help to visualize
 863 the distribution of the deviations in terms of sun eleva-
 864 tion and cloudiness condition. [Figure 8](#) shows the rMBD
 865 and rRMSD diagrams for the site-adapted Heliosat-4 and
 866 the CIM-McClear at the LE site. The corresponding dia-
 867 grams for the original Heliosat-4 and CIM-ESRA estimates
 868 are similar and are omitted for brevity. The comparison
 869 is intended to compare the deviations of locally adapted
 870 models that have a different nature and use different satel-
 871 lite inputs. Upon inspection of this figure, the metrics'
 872 patterns observed are different, and a better performance
 873 across all-sky conditions is evident for the CIM-McClear
 874 model. The rMBD of this model remains small for all the
 875 conditions (between $\simeq \pm 8\%$) and the higher deviations
 876 (rRMSD $\simeq 20\text{-}25\%$) occur around solar noon for partial
 877 cloudiness conditions (intermediate k_T values). The small
 878 overall underestimation bias observed in the model can be
 879 associated to clear-sky conditions and, to a minor extent,
 880 to cloudy conditions close to solar noon (see [Figure 8b](#)).
 881 When the Sun is low, the discriminated performance of
 882 both Heliosat-4 and CIMs are similar (low rMBD and
 883 rRMSD) for all cloudiness condition. However, when the
 884 Sun is high, the behavior changes, and errors associated to
 885 clouds are much higher for the Heliosat-4 method (see for
 886 instance, the $\simeq +40\%$ rMBD observed under this condi-
 887 tion for this model). In fact, the rRMSD for cloudy condi-
 888 tion ($k_T < 0.4$) at high Sun's elevation angle ($\cos\theta_z > 0.8$)
 889 are around 35-50% for Heliosat-4 and around 10-25% for
 890 the CIMs. This different behavior is related to the satel-
 891 lites' FOV: under this condition the optical path of the
 892 MSG satellite observation differs strongly from the optical
 893 path of the solar radiation through the atmosphere, which
 894 is not the case for the GOES-East satellite observation.
 895 In particular, midday clouds are worse perceived by the
 896 MSG satellite, resulting in higher overestimation (positive
 897 bias) and uncertainty (higher rRMSD) for the Heliosat-4.
 898 Further, as these higher deviations occur at midday when
 899 solar applications produce the most, they have an impor-
 900 tant impact in solar yield assessments. Based on this, it
 901 is clear that GOES-East satellite imagery, which has a
 902 smaller viewing angle for the region, is a better choice for
 903 local solar resource assessment.

904 The performance observed for the original Heliosat-4

905 GHI estimates is consistent with those found in the litera-
 906 ture about the Heliosat family ([Rigollier et al., 2004](#); [Eissa
 907 et al., 2012](#); [Ineichen, 2014](#); [Eissa et al., 2015a](#); [Qu et al.,
 908 2017](#)) and it is between the expected uncertainty range for
 909 satellite-based models ([Perez et al., 2013](#)). As an example,
 910 in [Ineichen \(2014\)](#) a long term (8-years) uncertainty evalu-
 911 ation of the Heliosat-3 model (among other satellite-based
 912 models) was reported over 18 high quality measurement
 913 sites in Europe. A negligible bias and an rRMSD of 20%
 914 was found for the hourly estimates. The evaluations re-
 915 ported in the literature use the single pixel approach, so
 916 the hourly uncertainty showed here is lower due to the
 917 spatial smoothing. Also, most evaluations do not report
 918 the uncertainty of site-adapted versions. As mentioned
 919 previously, the evaluation of the Heliosat-4 method (not
 920 site adapted) on a hourly basis is provided by the CAMS
 921 website for several sites across the world. Two of these
 922 sites are close to the region (Buenos Aires and Florianópolis)
 923 but not representative of the broader Pampa Húmeda
 924 area. Reported rMBD for these sites are between $-5\text{-}0\%$
 925 and the rRMSD values are between 25-28%. The biases
 926 reported in this work for the original Heliosat-4 method
 927 are in the same range, being the maximum overestimation
 928 $+11.1 \text{ Wh/m}^2$ or $+2.6\%$ (for the oceanic RO site) and
 929 the maximum underestimation of -14.4 Wh/m^2 or -3.1%
 930 (at the LU site). The average RMSD found here for the
 931 Pampa Húmeda region is significantly lower, but consid-
 932 ering spatial smoothing, and is of 80.2 Wh/m^2 or 17.9%.
 933 Site adaptation reduces slightly the RMSD to 75.2 Wh/m^2
 934 or 16.8%. The value of the site-adaptation procedure for
 935 the region is quantified in a 1.1% reduction in rRMSD,
 936 which is significant, but not enough to achieve the reduced
 937 rRMSDs of 12-13% of the CIMs presented here due to the
 938 lower GOES-East satellite FOV in the region.

939 The JPT and BDJPT empirical models have been ad-
 940 justed and evaluated for the same region in [Alonso-Suárez
 941 et al. \(2012\)](#), using GOES-East satellite images with the
 942 same spatial smoothing as here. In that work, 3 stations
 943 are used for models adjustment (LE, TT and LB) and 4
 944 stations are used for validation. None of the validation
 945 sites is any of the sites used in this work. At the hourly
 946 level, a small overestimating bias of $+1.4\%$ and $+1.1\%$
 947 was obtained for each model, respectively. In terms of
 948 RMSD and KSI, the overall results for the JPT model
 949 (rRMSD = 18.6% and KSI = 16 Wh/m^2) are similar
 950 to those of the original Heliosat-4 (being the former lo-
 951 cally adjusted) and the results for the BD-JPT model
 952 (rRMSD = 14.0% and KSI = 10 Wh/m^2) are slightly
 953 above those of the CIMs considered here. Further perfor-
 954 mance evaluations between these models (using the same
 955 data set) are required, in order to have a fair comparison
 956 between alternative models for this region. However, this
 957 is not under the scope of the present work.

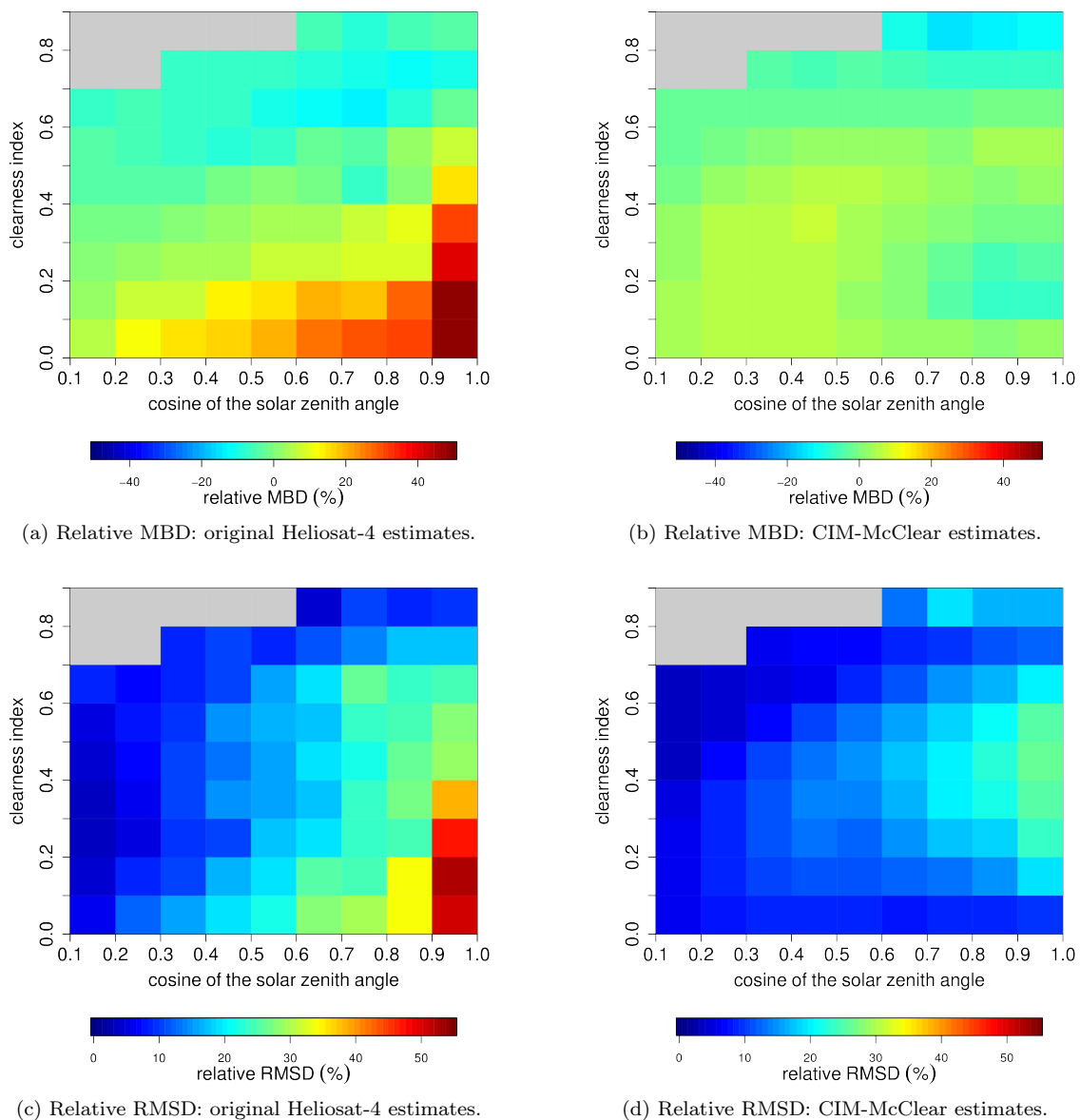


Figure 8: Performance metrics for the LE site discriminated by clearness index k_T and the cosine of the solar zenith angle for the site-adapted Heliosat-4 and the CIM-McClear estimates. The gray cells represent the absence of data for that condition. The values are relative to the measurements mean.

5. Conclusions

Different types of satellite-based models for estimating ground level solar global horizontal irradiation have been assessed at the hourly level using good quality ground data from 10 sites in the southeastern part of South America (Pampa Húmeda region). This region has simultaneous geostationary satellite coverage from the Meteosat Second Generation (MSG) and from the GOES-East satellites, with significantly different view angles over the area and thus provides an opportunity to quantify the effect of large view angles on the quality of the estimates.

The McClear and ESRA models have been considered for clear-sky estimation in the area. McClear estimates

have been compared to clear-sky ground data with and without site-adaptation. The ESRA clear-sky model has been implemented using a daily Linke turbidity cycle which captures the average seasonal trends of the local atmosphere (this model does not use satellite information). When compared to hourly clear-sky ground data both show a similar performance, with small biases and rRMSDs in the range $\approx 3\%$ - 4% of the measurement's average, which is similar to the ground measurement's uncertainty. McClear (even without site-adaptation) performs slightly better than the ESRA clear-sky model, due to its capacity to model the detailed atmospheric conditions on a daily basis. The site-adapted version of McClear performs best, resulting in unbiased estimates with 2.8% of rRMSD. However,

1085 the gain obtained from site-adaptation is small, 0.4% of
1086 rRMSD. To resolve between higher accurate clear-sky es-
1087 timates requires pyrhelimeter ground data accurate to 1%
1088 and is left for future work.

1089 The estimates from these clear-sky models have also
1090 been compared (one to the other) under all-sky conditions
1091 of the real atmosphere, which is uncommon in clear-sky
1092 models assessments. This comparison revealed a signifi-
1093 cantly different behavior of both models: when clouds are
1094 present in the real atmosphere, McClear estimates are sys-
1095 tematically lower than ESRA's. This is best quantified by
1096 the mean deviation between both estimates, which is 0.1%
1097 for clear-sky conditions and becomes -5.2% in the pres-
1098 ence of clouds (i.e. McClear estimates are about 5% lower
1099 than ESRA's). This different behaviour is due to the fact
1000 that McClear is sensitive to the short-term atmospheric
1001 variability (in particular, regarding water vapor content),
1002 while our implementation of ESRA is not. Since both are
1003 commonly used clear-sky models, this may be relevant in-
1004 formation for CIM-based all-sky models, short-term fore-
1005 casting, automated quality control or other applications
1006 that make use of a clear-sky index.

1007 For all-sky conditions, the Heliosat-4 (HS4) method
1008 and two CIMs (cloud index methods) have been consid-
1009 ered. A first representative assessment of the HS4 method
1010 for the Pampa Húmeda region is provided, with and with-
1011 out site-adaptation. This model is based on the McClear
1012 clear-sky model and MSG images. The two CIMs share the
1013 same formulation for the locally adapted cloud attenuation
1014 factor based on the cloud index derived from GOES-East
1015 satellite information. One of them (CIM-ESRA) is based
1016 on the ESRA clear-sky model and the other (CIM-McCclear)
1017 on the site-adapted McCclear model. The major difference
1018 in the satellite images used by both approaches is the satel-
1019 lite view angle over the region. Similar spatial smooth-
1020 ing and local-adaptation are applied to both, by different
1021 means.

1022 The HS4 estimates as provided by the CAMS-SoDa
1023 platform have the expected performance for a model with-
1024 out local adaptation but with spatial smoothing, showing
1025 small bias and rRMSDs in the range 16.6% - 20.4% (av-
1026 erage 17.9%) across stations. The site-adapted version of
1027 HS4 improves slightly this performance, showing no bias
1028 and rRMSD between 15.4% and 19.6% (average 16.8%).
1029 The overall gain in rRMSD due to site-adaptation is 1.1%
1030 under all-sky conditions. The implemented CIMs exhibit
1031 a small but consistent negative bias of about -1%, so a
1032 site-adaptation post-processing can be of practical rele-
1033 vance for these estimates. The rRMSD is in the range
1034 10.6% - 13.6% for CIM-McCclear and 11.2% - 13.6% for
1035 CIM-ESRA. CIM-McCclear has slightly smaller rRMSDs
1036 than CIM-ESRA, with averages of 12.1% for the former
1037 and 12.5% for the latter. These results for the locally ad-
1038 justed CIMs are comparable to the best results found in
1039 the literature and similar to those found for an empirical
1040 model optimized for Uruguay's territory (BD-JPT model).
1041 However, the empirical nature of this model implies that

1042 its generalization to all the Pampa Húmeda area or other
1043 areas of the continent is not straightforward, since its co-
1044 efficients present more spatial variability than the param-
1045 eters used for local adaptation in this work. These CIMs
1046 represent a significant improvement for satellite-based solar
1047 resource assessment over the extended region.

1048 This all-sky assessment implies that the satellite view
1049 angle over the area must be taken into account when es-
1050 timating ground level solar irradiation: relatively simple
1051 CIMs using lower view angle satellite information outper-
1052 form the sophisticated Heliosat-4 method which uses de-
1053 tailed atmospheric information and radiative transfer cal-
1054 culations. The MSG satellite views the region with a view
1055 angle of approximately 70°, while the GOES-East satel-
1056 lite has a viewing angle over this area of about 40°. It
1057 is shown that the impact of using satellite-based estima-
1058 tion out of the recommended area (satellite zenith angle
1059 larger than 60°) can easily account for the performance
1060 difference between MSG-based and GOES-based models
1061 observed over this region. For low solar altitude, both
1062 CIMs and HS4 present similar uncertainty for all cloudi-
1063 ness conditions. However, for high solar altitude, when
1064 the radiation optical path in the atmosphere is similar to
1065 that of the GOES-East but very different to that of the
1066 MSG, significantly higher errors can be observed for the
1067 HS4 model especially when partial cloudiness is present.
1068 Hence the performance difference is mostly explained by
1069 the satellites' cloud perception when the Sun is close to
1070 the zenith, which is directly related to each satellite FOV.
1071 The performance difference presented here should be read
1072 as an example of impact assessment of using solar satel-
1073 lite estimates out of the recommended area and is not a
1074 statement about the relative quality of the models.

1075 In sum, the site adapted McCclear clear-sky model is
1076 highly accurate in this region, but caution should be taken
1077 when using its estimates under all-sky conditions because
1078 it is sensitive to changes in atmospheric conditions under
1079 cloudy conditions. The ESRA model with local average T_L
1080 trends also gives good results for the region and it is insen-
1081 sitive to the presence of cloudiness in the real atmosphere.
1082 For all-sky estimates, it is not recommended to use MSG-
1083 based models over this area (even with site adaptation),
1084 due to the higher view angle and the associated decrease in
1085 accuracy. Both CIMs, based on GOES-East satellite im-
1086 ages show a remarkable performance over this region, pro-
1087 vide accurate hourly estimates for global solar irradiance
1088 and have the potential to be extended to a broader area.
1089 Furthermore, they can potentially be adapted to provide
1090 DNI estimates or, combined with spectral clear-sky mod-
1091 els, provide spectral estimates for global irradiation.

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